

Figure 35. Location map of principal localities in the Wichita Mountains showing Rowe Quarry, where samples were collected. (After Powell and others, 1980.)

alytical data were interpreted by fitting a line to the data points, using the least-squares-cubic method of York (1966), and solving for the intersection with concordia (Wetherill, 1956).

Nine magnetic fractions of zircon from the Mount Sheridan Gabbro and its associated pegmatitic pods were analyzed. The zircons from both the gabbro and pegmatite are similar in morphology, being mostly clear and euhedral and relatively free of inclusions (fig. 37). Zircon was much more abundant (approximately 200 times) in the pegmatite than in the gabbro.

The U/Pb data from the individual fractions (table 16) define separate chords on a U/Pb concordia diagram (fig. 38). The chord defined by data from the pegmatite intersects concordia at  $552 \pm 7$  m.y. This age is believed to represent the time of crystallization of the Mount Sheridan Gabbro pegmatite. The chord

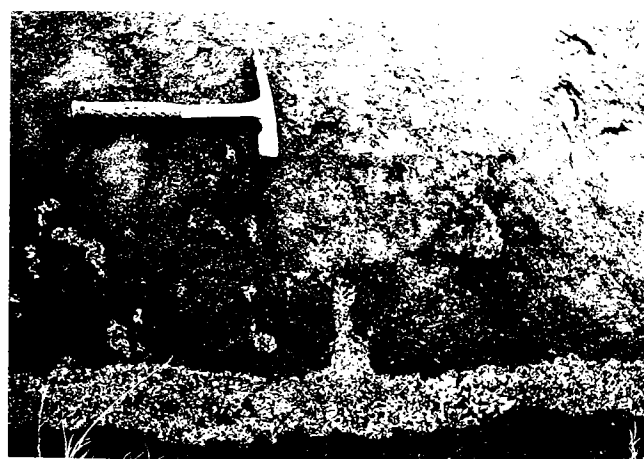


Figure 36. Pegmatitic pods in Mount Sheridan Gabbro, Rowe Quarry. (Photograph courtesy of M. C. Gilbert.)